

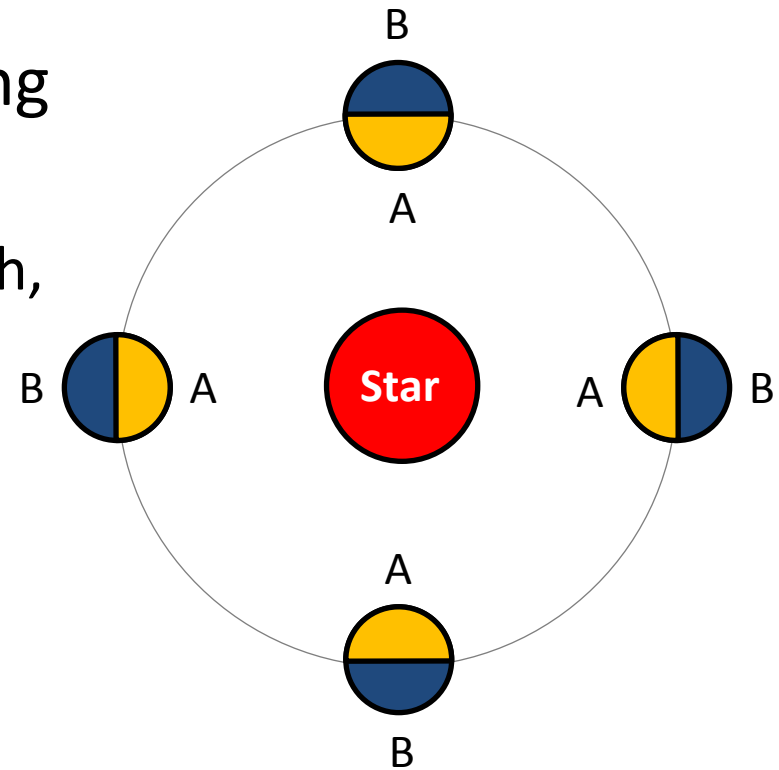
A numerical study on atmospheric general circulations of synchronously rotating aqua-planets: Dependence on planetary rotation rate and Solar Constant

S. Noda^(1,6), M. Ishiwatari^(2,3), K. Nakajima⁽⁴⁾, Y. O. Takahashi^(1,3),
S. Nishizawa⁽⁷⁾, M. Onisi^(1,3), G. L. Hashimoto⁽⁵⁾,
K. Kuramoto^(2,3), Y.-Y. Hayashi^(1,3)

(1) Kobe University, Kobe, (2) Hokkaido University, Sapporo,
(3) Center for Planetary Science, Kobe, (4) Kyushu University, Fukuoka,
(5) Okayama University, Okayama, (6) Meteorological Research Institute, Tsukuba,
(7) RIKEN Advanced Institute for Computational Science, Kobe

Synchronously rotating planets

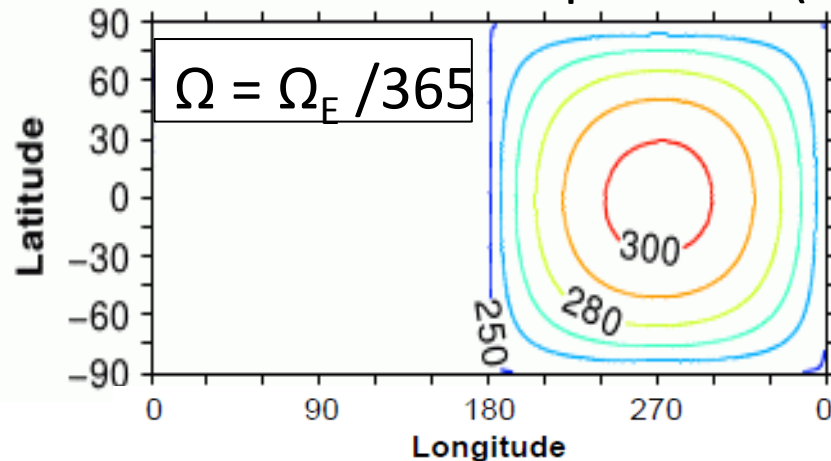
- Some terrestrial exoplanets are expected to be tidally locked and synchronously rotating.
 - fixed dayside and nightside
- Climates of synchronously rotating aqua-planets?
 - With solar constant of present Earth, day-side is in runaway greenhouse condition



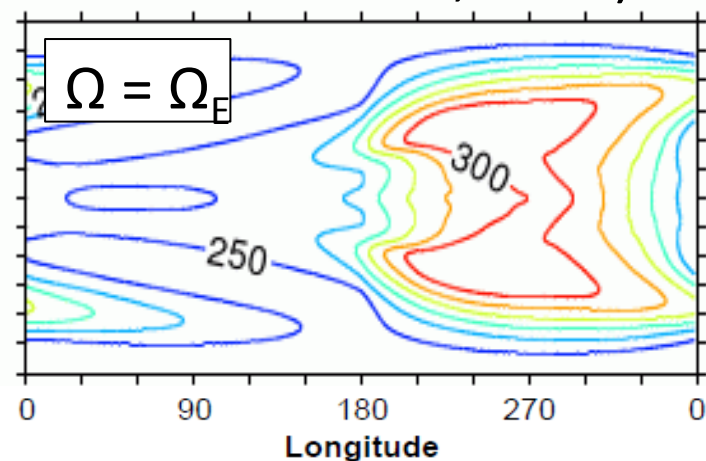
Previous aqua-planet experiments

- GCM experiment with changing rotation rate.
 - Joshi (2003): $\Omega/\Omega_E=1$
 - Merlis and Schneider (2010): $\Omega/\Omega_E=1/365 \sim 1$
 - Edson et al. (2011): $\Omega/\Omega_E=1/100 \sim 1$
- Equilibrium states are obtained in all cases.

Surface temperature (Merlis and Schneider, 2010)



Slowly rotating regime



Rapidly rotating regime

Aims of this study

- Parameter study with various planetary rotation rate (Ω) and solar constant (S)
 - Parameter dependences of circulation and energy transports ?
 - In previous works, cases with $\Omega \sim 0, \Omega_E$ are focused on.
 - Occurrence condition of the runaway greenhouse state?
 - The dependence on solar constant has not been examined in previous works.

Model

- Atmospheric General Circulation model : dcpam (<http://www.gfd-dennou.org/library/dcpam/index.htm.en>)
- Dynamical part : 3-d primitive equation on sphere

Hydrostatic equation

$$\frac{\partial \Phi}{\partial \sigma} = -\frac{R^d T_v}{\sigma}$$

Equation of motion

$$\frac{du}{dt} - fv - \frac{uv}{a} \tan \varphi = -\frac{1}{a \cos \varphi} \frac{\partial \Phi}{\partial \lambda} - \frac{R^d T_v}{a \cos \varphi} \frac{\partial \pi}{\partial \lambda} + \mathcal{F}_\lambda,$$

$$\frac{dv}{dt} + fu + \frac{u^2}{a} \tan \varphi = -\frac{1}{a} \frac{\partial \Phi}{\partial \varphi} - \frac{R^d T_v}{a} \frac{\partial \pi}{\partial \varphi} + \mathcal{F}_\varphi.$$

Continuity equation

$$\frac{d\pi}{dt} + \nabla \cdot \mathbf{v}_H + \frac{\partial \dot{\sigma}}{\partial \sigma} = 0. \quad \mathbf{v}_H \cdot \nabla_\sigma = \frac{u}{a \cos \varphi} \frac{\partial}{\partial \lambda} + \frac{v}{a} \frac{\partial}{\partial \varphi}$$

Equation of heat

$$\frac{dT}{dt} = \frac{R^d T_v}{C_p^d} \left\{ \frac{\partial \pi}{\partial t} + \mathbf{v}_H \cdot \nabla_\sigma \pi + \frac{\dot{\sigma}}{\sigma} \right\} + \frac{Q^*}{C_p^d}.$$

Conservation of water vapor

$$\frac{dq}{dt} = S_q$$

Boundary Condition

$$\Phi = 0 \quad \text{at } \sigma = 1$$

$$\dot{\sigma} = 0 \quad \text{at } \sigma = 0,1$$

(λ, φ): longitude, latitude

$\pi = \ln p_s$

$\sigma = p / p_s$

$\dot{\sigma} = d\sigma/dt$

p : pressure

p_s : surface pressure

a : planetary radius

u : zonal wind

v : meridional wind

\mathbf{v}_H : horizontal velocity

T : temperature

T_v : virtual temperature

Φ : geopotential

f : Coriolis parameter

R^d : gas constant of dry air

C_p^d : specific heat of dry air

q : specific humidity

Q^* : external Heating

S_q : Source of water vapor

F_λ, F_φ : external forcin

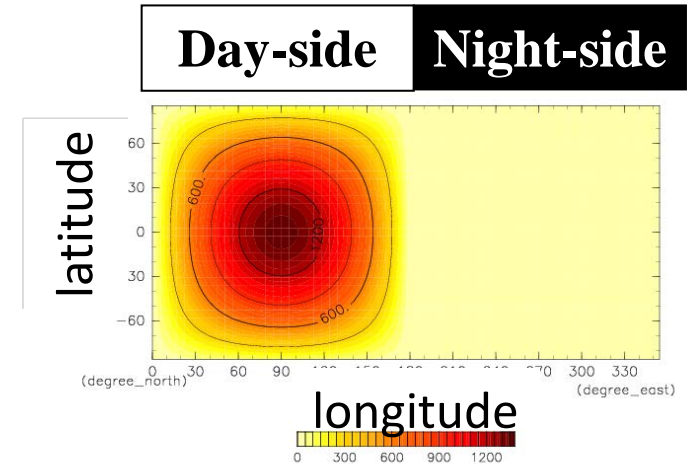
Physical Processes

Atmospheric composition	“Water vapor” and “dry gas”
Radiation process	Water vapor : gray to IR radiation Dry gas: transparent
Cumulus process	Convective adjustment (Manabe et al., 1965) condensed water is removed from system immediately (no cloud)
Turbulent vertical mixing	Mellor and Yamada (1972) Level 2
Surface condition	Aqua-planet, flat surface Zero heat capacity, no sea ice

Experimental setup

- Radius of planets, mean surface pressure, gravitational acceleration, and so on are Earth's values
- Obliquity is 0

Incoming radiation flux



	Ω dependence exp.	S dependence exp.
Rotation rate Ω (Ω_E : Earth's value)	0 - Ω_E (16 cases)	0, 0.15 Ω_E , 0.5 Ω_E , Ω_E
Solar constant S (global mean flux)	1380 W/m ² (345 W/2)	1380 - 1700 W/m ² (24 cases) (345 - 425 W/m ²)
Resolution	T21L16 (64 x 32 x 16)	T21L32 (64 x 32 x 32)
Integration time	2000 day (last 1000 days are analyzed)	Over 2000 days (last 500 days are analyzed)
Time step	20 min	10 min or less

Dependence on rotation rate

- Experiment
 - Various Ω , fixed solar constant
- Results
 - In all cases, the runaway greenhouse states do not occur: statistically equilibrium state or oscillating state
 - Day-night energy transport is independent of Ω . Circulation pattern depends on Ω .

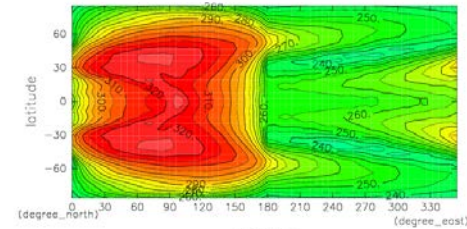
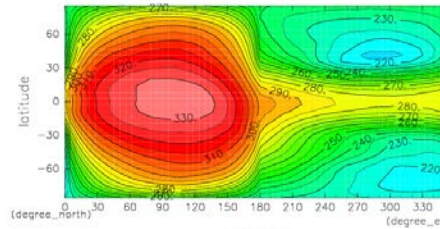
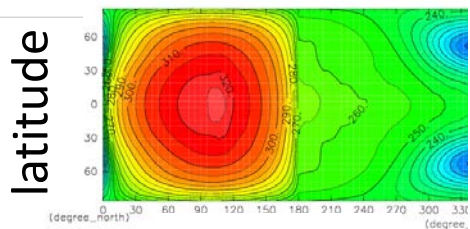
Atmospheric structure for various Ω

$\Omega = 0.05 \Omega_E$

$\Omega = 0.67 \Omega_E$

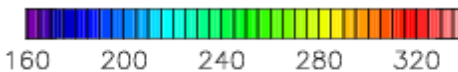
$\Omega = \Omega_E$

Surface Temperature

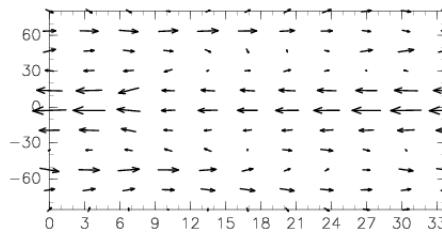
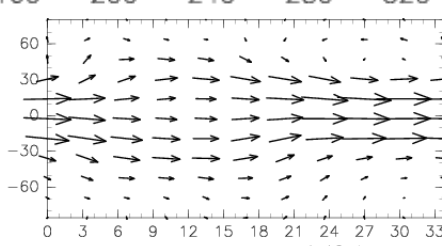
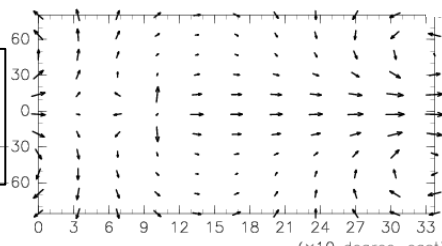


latitude

longitude



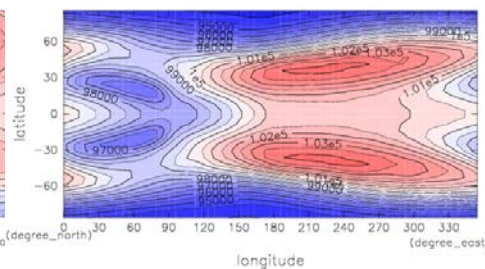
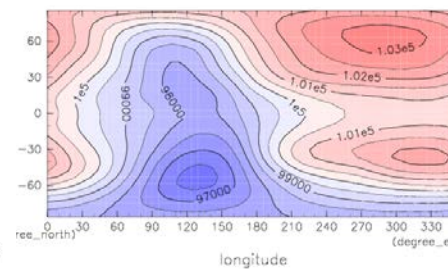
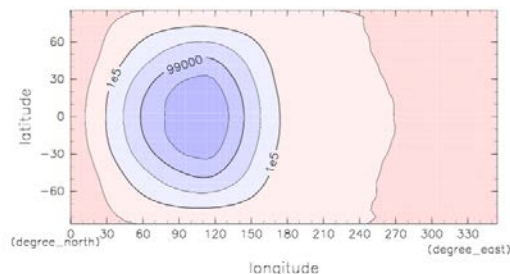
Horizontal velocity at sigma=0.2



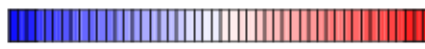
100 m/s

↑
→ 100 m/s

Surface pressure



93000 Pa



107000 Pa

Time mean over 1000-2000day

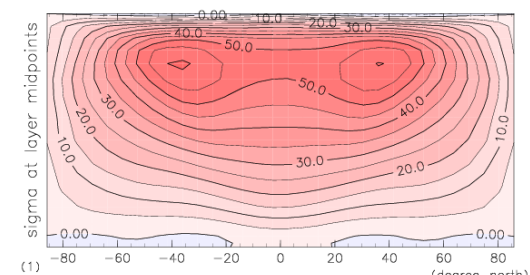
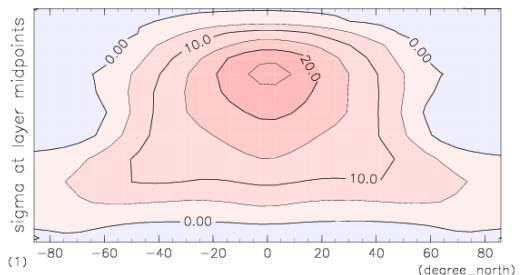
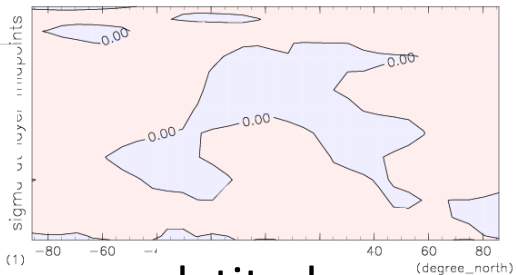
Zonal mean zonal wind for various Ω

$\Omega = 0$

$\Omega = 0.05 \Omega_E$

$\Omega = 0.15 \Omega_E$

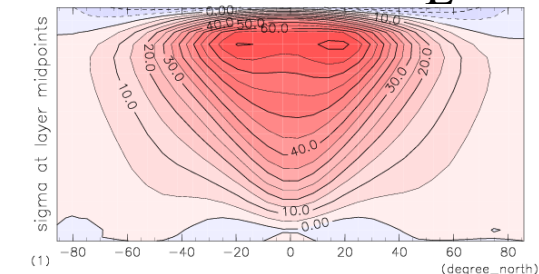
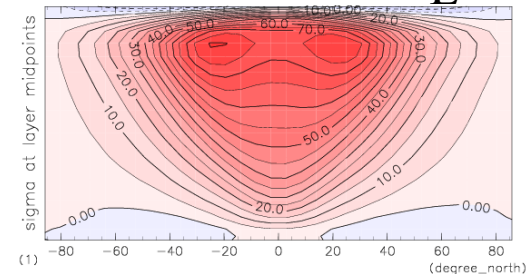
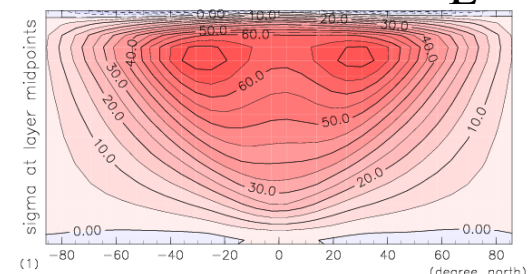
sigma



latitude
 $\Omega = 0.25 \Omega_E$

latitude
 $\Omega = 0.33 \Omega_E$

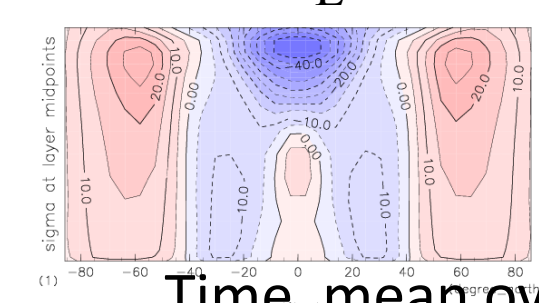
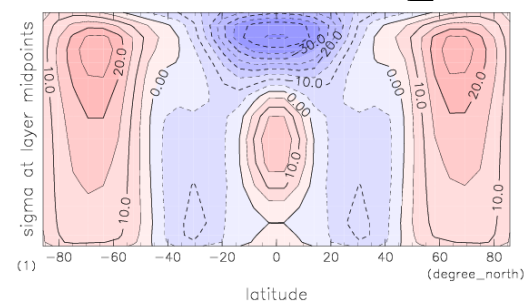
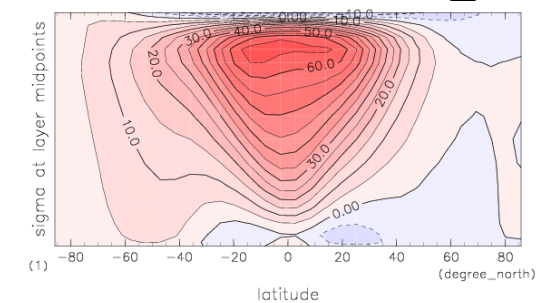
latitude
 $\Omega = 0.5 \Omega_E$



latitude
 $\Omega = 0.67 \Omega_E$

latitude
 $\Omega = 0.8 \Omega_E$

latitude
 $\Omega = \Omega_E$



CONTOUR INTERVAL = 5.000E+00

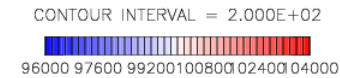
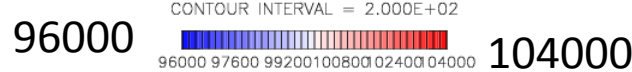
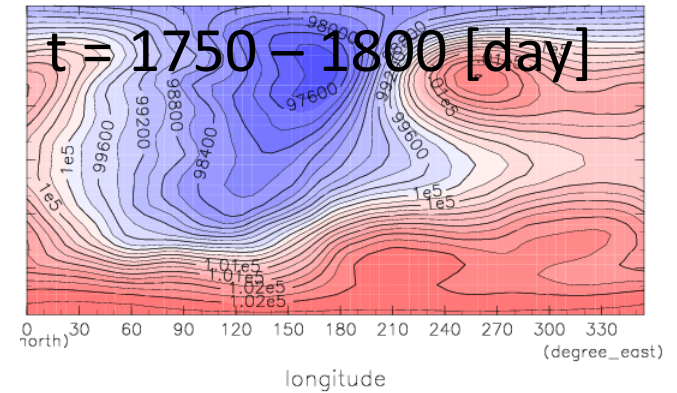
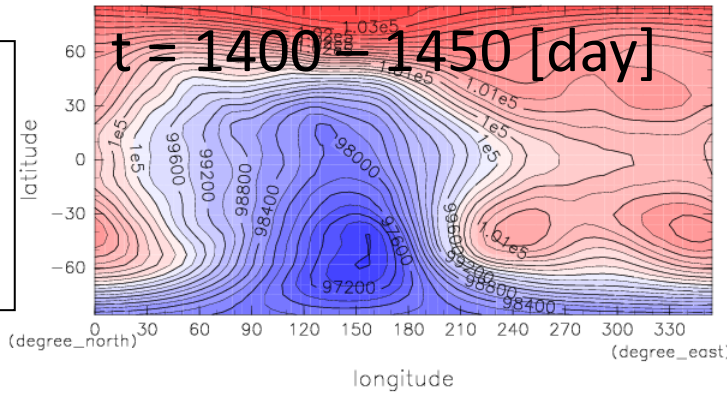


Time mean over
1000-2000day

Asymmetric state: $0.2\Omega_E - 0.67\Omega_E$

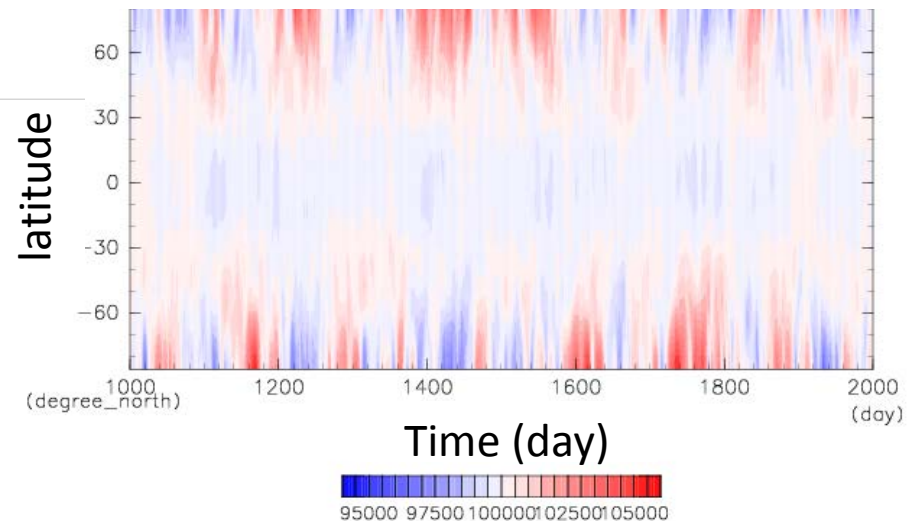
Surface pressure for $0.5\Omega_E$

North-south
Asymmetric
State

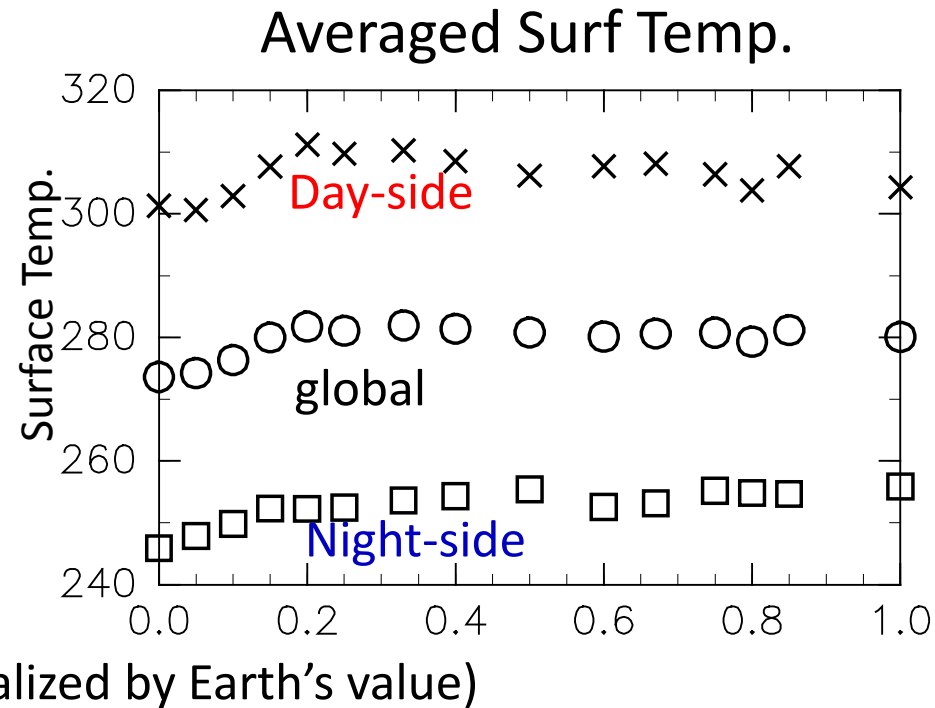
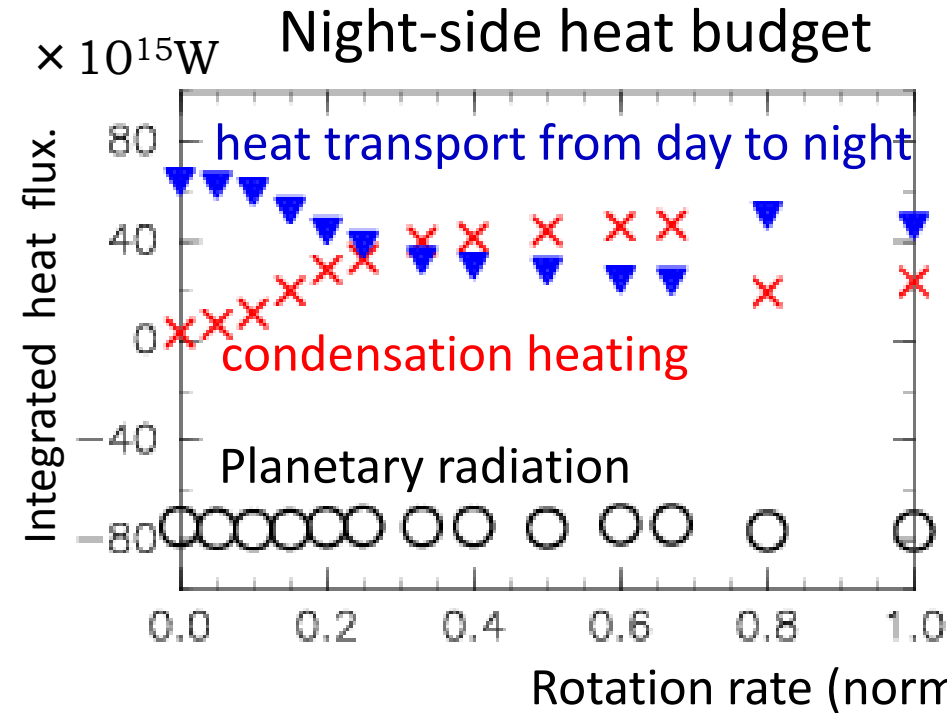


- Oscillation
 - Non-periodic for large Ω
 - “period”: 10day – 1000day
 - also in high reso. experiments

zonal mean surface pressure



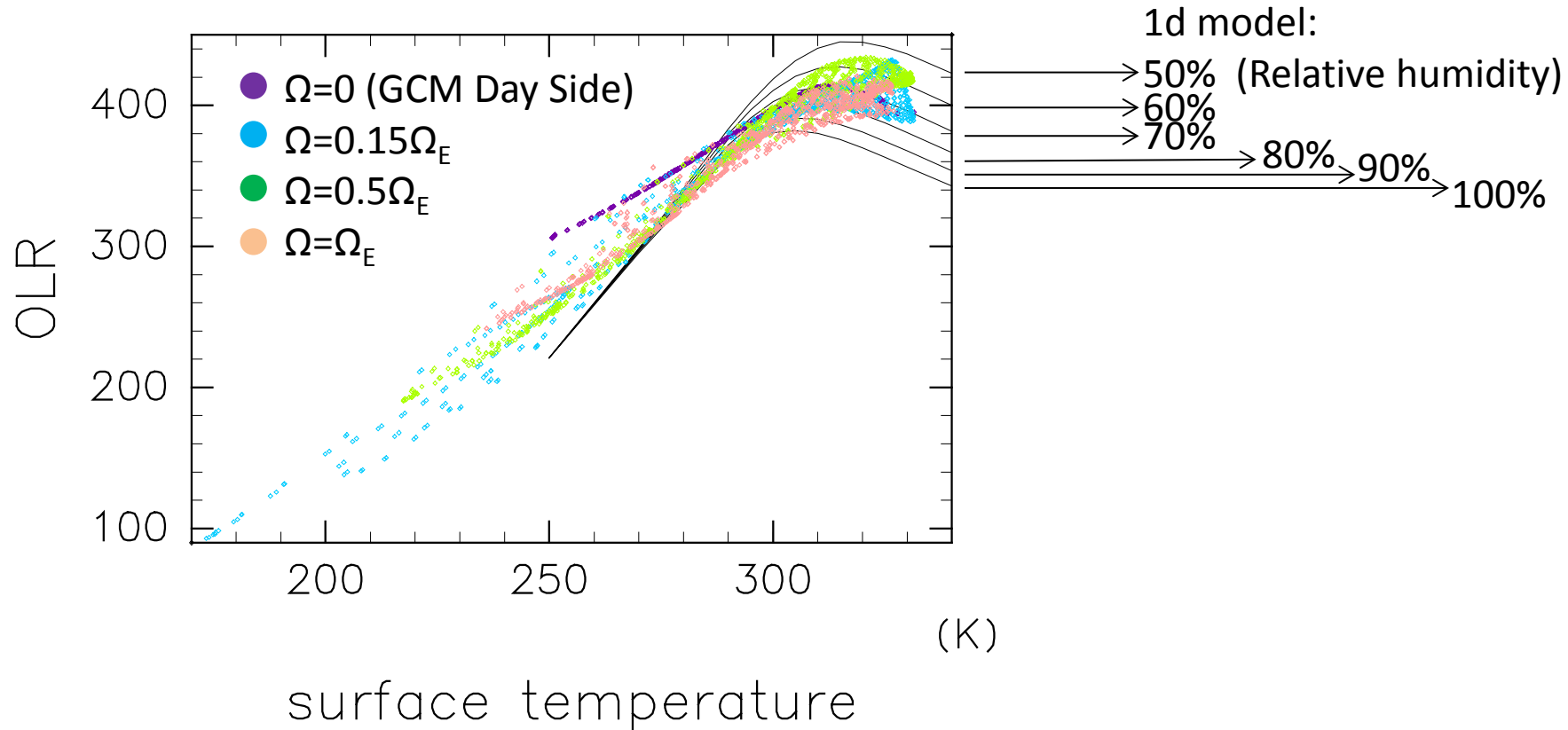
Dependence of heat budget on Ω



- Small dependence of summation of sensible/latent heat transports on Ω
- Total heat transport may be determined by (Incident solar flux) – (radiation limit)
 - Radiation limit: Nakajima et al. (1992), Ishiwatari et al. (2002)

Comparison with 1-d model

(W m⁻²)



- Outgoing Longwave Radiation does not exceed radiation limit obtained by 1-dim radiative convective equilibrium model.

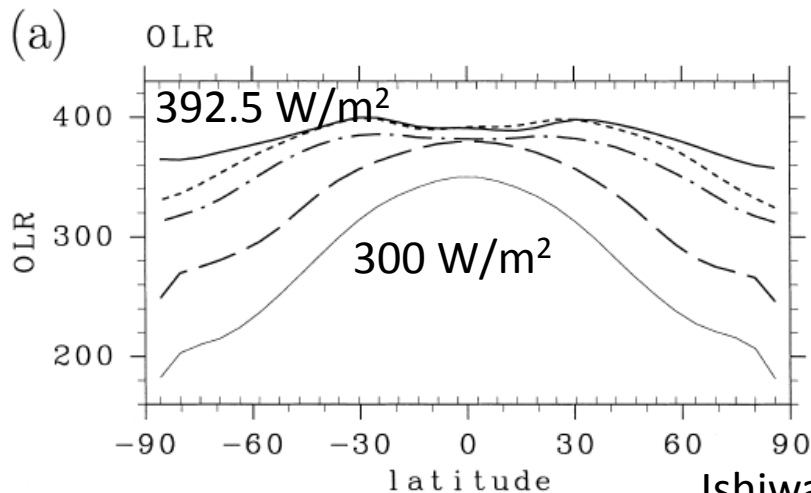
Dependence on solar constant

- Experiment

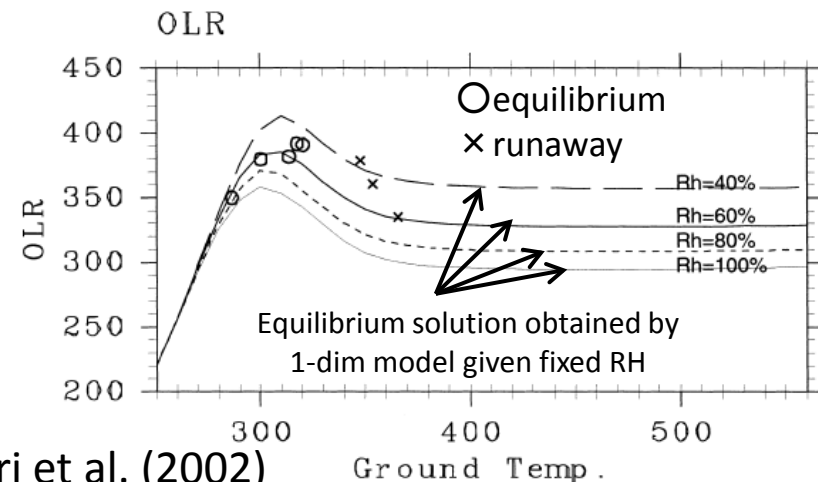
- Various solar constant, $\Omega=0$, $0.15 \Omega_E$, $0.5 \Omega_E$, Ω_E

- Aims

- Does runaway condition corresponds to the radiation limit obtained by 1-dim model?
 - In non-synchronous rotating cases, runaway condition corresponds to the radiation limit.

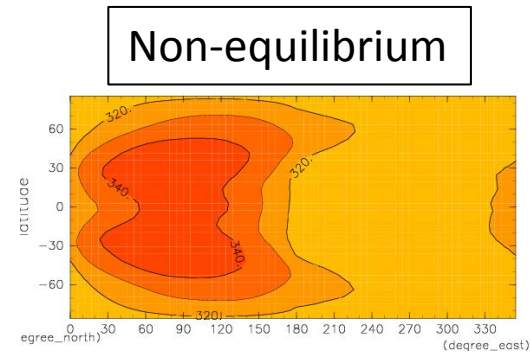
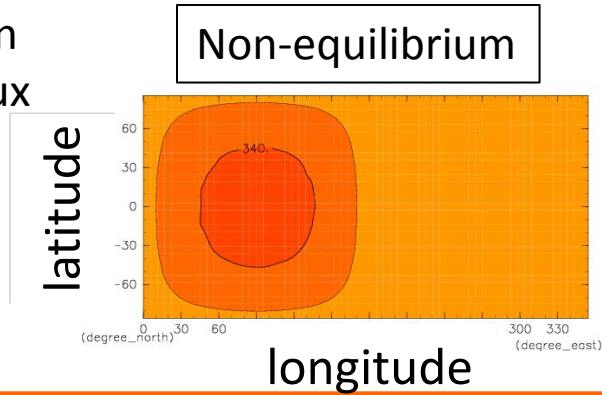


Ishiwatari et al. (2002)

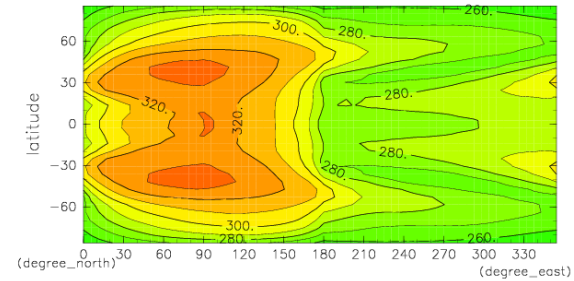
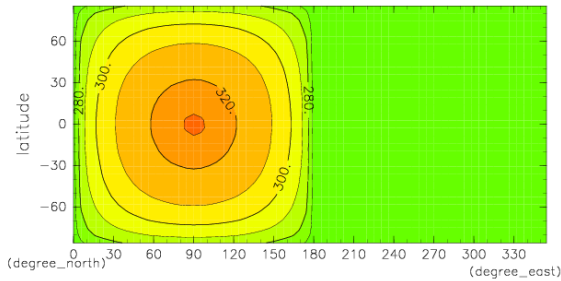


Occurrence of Runaway greenhouse states

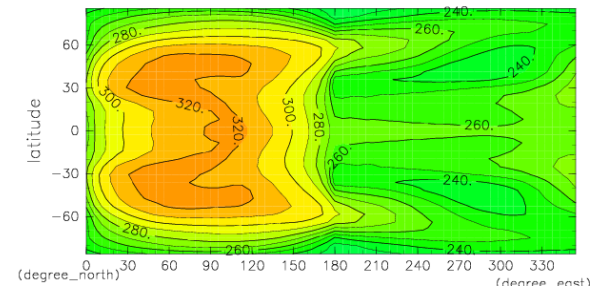
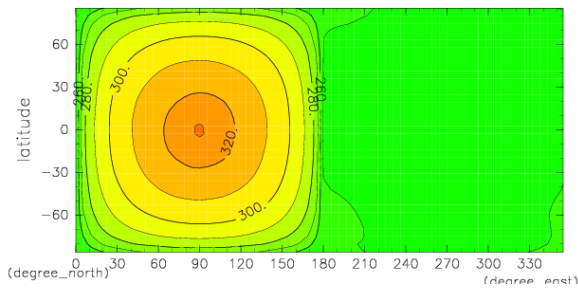
Global mean incoming flux
(W/m^2)
425



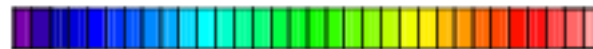
375



345



0

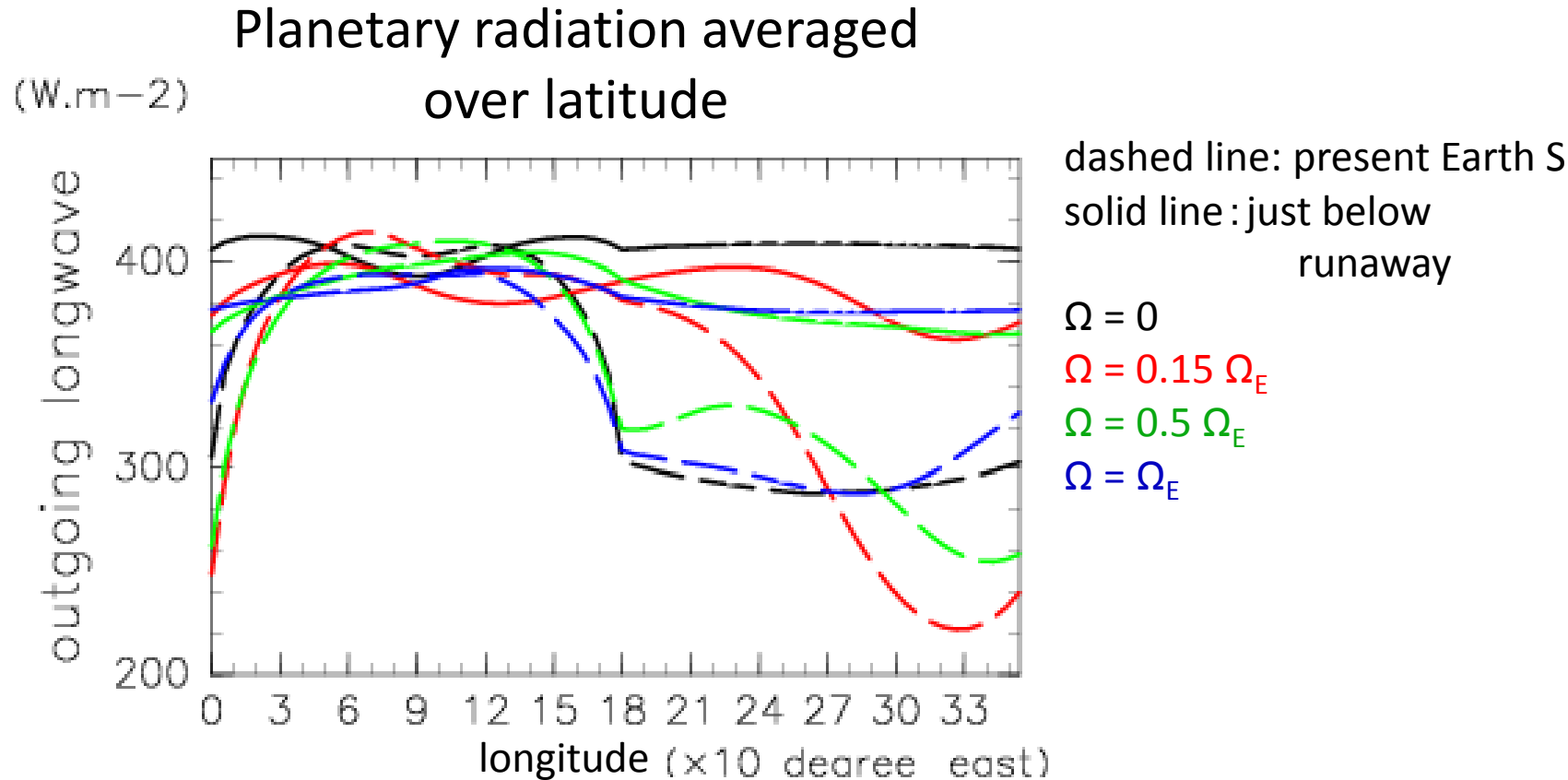


140 210 280 350

Ω_E

rotation
rate

OLR difference between day and night



- day-side OLR: almost constant against increasing S
- night-side OLR: increase with increasing S

Summary

- Amount of day-night energy transport is independent of planetary rotation rate.
 - Day-side planetary radiation is constrained by radiation limit obtained by vertical 1-d radiative-convective model.
- It seems that occurrence condition of the runaway greenhouse state is described by the radiation limit.
 - Difference of planetary radiation between day-side and night-side decreases with increased solar constant.
- Circulation patterns does not depend on solar constant but on planetary rotation rate.
 - Day-night direct circulation, super rotation, oscillating asymmetric states